

PHY-302

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Nuclear Decay

What is nuclear decay? <u>Spontaneous</u> transition from one state to another $X \Rightarrow Y + ...$



We will return to these in much more detail in future lectures. For now we concentrate on properties common to all types of decay



Radiation Stopping Distances

Each radiation type has its own degree of penetration in different materials



 β Radiation typically 3 m of air few centimeters water thin layer of glass / metal

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Decays are statistical: cannot predict when any particular nucleus will decay

Process follows Poisson Statistics



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Multimodal Decays



Unstable nuclei can often decay via 2 or more modes e.g. α and β decay modes Such nuclei have multimodal decays

Each decay mode is random and independent of the other modes

Each mode has it's own transition probability i.e. it's own λ or τ

$$\frac{\mathrm{d}N}{\mathrm{d}t} = \lambda_1 N - \lambda_2 N \qquad \longrightarrow \qquad N(t) = N_0 e^{-(\lambda_1 + \lambda_2)t}$$

The total decay constant is the sum of partial decay constants

$$\lambda = \lambda_1 + \lambda_2$$
 or $\lambda = \sum \lambda_i$

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Radioactive Production and Sequential Decay

Radioactive material often produced by exposure to neutron flux Neutron absorbed by nucleus - often followed by β decay e.g.

$$\begin{array}{rcl} & {}^{23}_{11}\mathrm{Na}+n & \rightarrow & {}^{24}_{11}\mathrm{Na} & \rightarrow & {}^{24}_{12}\mathrm{Mg}+e^-+\bar{\nu}_e \end{array} \qquad & \text{Rate of change of 24Na nuclei} \\ & \text{Differential equation for change in number of 24Na nuclei:} \\ & \text{where } p \text{ is production rate of 23Na (contributes with + sign)} \qquad & \underline{\mathrm{d}N} \\ & \underline{\mathrm{d}N} = p - \lambda N \end{array}$$

Materials often have decay chains: parent produces daughter products which are also radioactive

For a decay chain $N_1 \rightarrow N_2 \rightarrow N_3$ production/decay rates are given by

Differential

$$\frac{\mathrm{d}N_1}{\mathrm{d}t} = -\lambda_1 N_1$$
$$\frac{\mathrm{d}N_2}{\mathrm{d}t} = -\lambda_2 N_2 + \lambda_1 N_1$$



For chain of decay reactions: $N_1 \rightarrow N_2 \rightarrow N_3$

then
$$dN_2 = \lambda_1 N_1 dt - \lambda_2 N_2 dt$$

try solution: $N_2(t) =$

$$N_2(t) = Ae^{-\lambda_1 t} + Be^{-\lambda_2 t}$$

with boundary conditions

$$\begin{split} N_1(t=0) &= N_0 & \underset{\lambda_2(t)=N_0}{\overset{N_2(t)=N_0}{\frac{\lambda_1}{\lambda_2-\lambda}(e)}} \\ N_1(t=0) &= 0 & \text{Initial sample is pure } {}^{234}\text{U} \\ N_2(t=0) &= 0 & \text{No } {}^{230}\text{Th at time t=0} \\ N_3(t=0) &= 0 & \text{No } {}^{226}\text{Ra at time t=0} \end{split}$$

thus
$$N_2(t)=N_0rac{\lambda_1}{\lambda_2-\lambda_1}(e^{-\lambda_1 t}-e^{-\lambda_2 t})$$

hint: maximum production is reached when rate of change = 0

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Radioactive Decay	

- Decay Law is a probability relation need to take into account possible statistical fluctuation in measurements
- Decay time is independent of real time at t=0, i.e. cannot tell age of sample

Activity A , number of decays per second $~A=-\frac{\mathrm{d}N}{\mathrm{d}t}$ $=-\lambda N(t)$ $=\lambda N_0 e^{-\lambda t}$

How radioactive is a sample?:

Becquerel (Bq) : modern SI unit A radioactive sample with an activity of I Bq has an average of I decay per s.

Activity of a sample is affected by how much you have AND how old it is more material = more activity older sample = less activity

Other units attempt to quantify the hazard posed by radioactive sources e.g. Sieverts, Grays - will define these later in discussion of medical applications



Natural Radioactivity

Earth formed ~ 4.5 billion years ago

Material rich in Fe, C, O, Si

Heavier elements created in process of Nucleosynthesis from supernovae will be covered in more detail later

Only few elements have half-lives similar to age of earth

Element	Nucleus	Half-life /years
Thorium	²³² Th	1.41 x 10 ¹⁰
Neptunium	²³⁷ Np	2.14 × 10 ⁶
Uranium	²³⁸ U	4.47 × 10 ⁹
Actinium	²³⁵ U	7.04 × 10 ⁸

These are not the only sources of natural radioactivity shorter half-life sources continuously formed on earth today $^{14}\rm{C}$ is a useful example

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Organic matter absorbs CO₂ from atmosphere

In air stable ¹²C is most abundant (98.89%)

¹³C is remaining 1.11%

Radioactive ¹⁴C exists in 1 atom in 10¹² of ¹²C

Produced in upper atmosphere by cosmic rays:

 ^{14}N + n \rightarrow ^{14}C + p

Half-life is 5730y

When organism dies it stops acquiring $^{14}\mbox{C}\,$ - no longer in equilibrium state Remaining $^{14}\mbox{C}\,$ decays as usual

Age of samples determined by measuring specific activity (activity per gram)

Activity
$$A = -\frac{\mathrm{d}N}{\mathrm{d}t}$$
 (rate at which decays occur)
= $-\lambda N(t)$
= $\lambda N_0 e^{-\lambda t}$

Can also use mass spectrometers to determine ¹⁴C ratios

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A sample of wood used in constructing an ancient shelter yields 2.1 ¹⁴C decays per minute A similar sample of wood cut recently from a tree yields 5.3 decays per minute How old is the archeological sample?

$$N = N_0 e^{-\lambda t} \qquad \qquad t_{1/2} = \frac{\ln 2}{\lambda} = 5730 \text{ years}$$

consider the activity of sample when it died (t=0) and now (t=T)

$$\frac{\Delta N}{\Delta t}|_{t=0} = A_0 = \lambda N_0 \qquad \text{ and } \qquad \frac{\Delta N}{\Delta t}|_{t=T} = A_T = \lambda N(T) = \lambda N_0 e^{-\lambda T}$$

organism comes out of

organism comes out of $e^{\lambda T} = \frac{A_0}{A_T} \longrightarrow T = \frac{1}{\lambda} \ln \frac{A_0}{A_T}$ $= \left(\frac{t_{1/2}}{\ln 2}\right) \ln \frac{A_0}{A_T}$ = 7653 years

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Radio-carbon Dating

Technique assumes constant production of ¹⁴C over 50,000 years

tested by counting tree rings and historical records

not used for samples more than \sim 50,000 years old; count rate too low

can use other decay chains in similar manner - different half-life / assumptions

However, in last century fossil fuel burning increases stable carbons in air Also atmospheric atomic bomb tests increase ¹⁴C in air

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